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# Spatial soil loss estimation using an integrated GIS-based revised universal soil loss equation (RUSLE) in selected watersheds in northern Ghana

T. A. Adongo<sup>a,\*</sup>, W. A. Agyare<sup>b</sup>, F. K. Abagale<sup>c</sup>, N. Kyei-Baffour<sup>b</sup>

<sup>la</sup>Department of Agricultural Mechanisation & Irrigation Technology, University for Development Studies, Tamale, GHANA <sup>b</sup>Department of Agricultural and Biosystems Engineering, Kwame Nkrumah University of Science and Technology, Kumasi, GHANA <sup>c</sup>Department of Agricultural Engineering, University for Development Studies, Tamale, GHANA Corresponding Author: adongo.apusiga@yahoo.com

### Abstract

This study used thespatially integrated RUSLE to estimate annual soil loss and spatial distribution of soil loss severity classes for nine (9) reservoir catchments in northern Ghana. This model is an empirical model with five (5)main input parameters: rainfall erosivity, soil erodibility, slope length/steepness, landcover management and erosion management practice factors. Estimated annual soil loss rate ranged from 0 - 96.30 t/ha/y. Estimated mean annual soil loss ranged from 3.71 - 8.17 %. The severity of annual soil loss rates ranged from very low class (0.0 - 1.0 t/ha/y) to very high class (> 60.0 t/ha/y). Across the nine (9) catchments, the very low soil loss severity class was noted to constitute a larger portion of between 36.70 to 67.50 %. The moderate soil severity class was noted as the highest contributor to total annual soil loss. Soil and water conservation measures are required in the watershed to reduce soil loss.

Keywords: Watershed, soil loss, Integrated GIS-based RUSLE model, surface water resources, lixisols

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#### 1. Introduction

Soil loss in watersheds of surface water resources is a serious environmental problem (Gelagay and Minale, 2016) with the root cause being erosion (Bai *et al.*, 2008). Apart from causing accelerated on-site soil nutrient loss from farmlands in affected watersheds, it has resulted in accelerated off-site sediment accumulation in reservoirs, with implications for severe reduction in the designed life and water storage capacities of the affected reservoirs (Wang *et al.*, 2018). Though soil loss is a natural geological phenomenon and the result of the interplay between rainfall erosivity and soil erodibility factors, inappropriate human practices such as deforestation, cultivation in upslope areas without any support practices, bush burning, extension of urban areas, and uncontrolled and overgrazing have significantly aggravated soil loss in watersheds worldwide (Shiferaw, 2011;Mekonnen*et al.*, 2017).

In Ghana, Amegashie (2009) reported high soil loss rates of 18.28 – 157.55 t/ha/y in five (5) watersheds in the Upper East Region. Mean soil loss rates estimated in some watersheds of semi-arid areas of Africa include 22.80 t/ha/y at Tono, Ghana (Abubakari, 2014), and 47.4 t/ha/y at Koga, Ethiopia (Gelagay and Minale, 2016). Assessment of soil loss is useful in planning and soil conservation works in watersheds (Ganasri and Ramesh, 2015) with substantial efforts made on the development of soil loss assessment models (Nearing *et al.*, 2005) to quantitatively assess the extent and magnitude of soil loss (Kothyari*et al.*, 1994).

Soil loss information per unit land area in a watershed can be assessed using several models such as the old Universal Soil Loss Equation (USLE) - Empirical model (Wischmeier and Smith, 1978), Revised Universal Soil Loss Equation (RUSLE) - Empirical model (Renardet al. 1997), European Soil Erosion Model (EUROSEM) - Physically-based model (Morgan et al., 1998), SystemeHydrologiqueEuropeen or European Hydrological System (MIKE-SHE) - Physically-based model (Abbott et al., 1986), Water Erosion Prediction Project (WEPP) - Physically-based model (Laflenet al., 1991), Soil and Water Assessment Tool

(SWAT) - Conceptual model (Arnold et al., 1993) and Agricultural Catchment Research Unit (ACRU) - Conceptual model (Schulze, 1995).

Empirical models are a simulation of natural processes, mostly based on statistical observations and rely on developed regression relationships. The computational processes of empirical models are simple and their data requirements are less than those that are required for conceptual and physically-based models (Wheater*et al.*, 1993; Hajigholizadeh*et al.*, 2018). Physically-based models are generally based on the concept of the conservation of mass, momentum equations and energy as governing equations describing streamflow or overland flow, and conservation of mass equation for sediment (Bennett, 1974; Kandel*et al.*, 2004). Conceptual models are basically a combination of empirical and physically-based models and are more applicable to answering general questions (Beck, 1987; Hajigholizadeh*et al.*, 2018). These models were developed on the basis of spatially-lumped forms of water and the sediment continuity equation. The main focus of a conceptual model is to predict sediment yield, basically using the concept of the unit hydrograph (Lal, 1994).

Using conventional methods, physically-based models or conceptual models to assess soil loss is very expensive and time consuming so a RUSLE model integrated with remote sensing and GIS was used for the study whilst noting the non-existence of soil loss information (Wheater*et al.*, 1993; Hajigholizadeh*et al.*, 2018). The Integrated GIS-based RUSLE model was also selected for its minimal number of required data and for its ease as a tool for field application and its ability to analyze soil loss potential on a cell-bycell basis (Shinde*et al.*, 2010).Thisstudy assessed the annual soil loss per unit area of land innine (9) watersheds in northern Ghana and the estimated soil loss ratesfor some watersheds were within the tolerable limits as indicated by the FAO (1984) and USDA-NRCS (1999.

#### 2. Materials and methods

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#### 2.1 Description of Study Areas

The study was carried out in nine (9) watersheds across the former three (3) administrative regions namely; Bontanga, Libga and Golingawatersheds in the Northern Region; Vea, Tono and Gambibgowatersheds in the Upper East Region and Sankana, Karni and Daffiamawatersheds in the Upper West Region of Ghana. A map showing the study watersheds is presented in Figure 1 whilst the description of the watersheds is presented in Table 1.

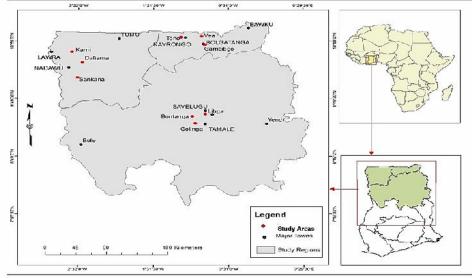


Figure 1: Map of the Former Three(3) Northern Regions Showing the Study Watersheds

Name of wa	tershed	Bontanga	Golinga	Libga	Gambibgo	Tono	Vea	Daffiama	Karni	Sankana
Location	Region	Northern	Norther n	Northern	Upper East	Upper East	Upper East	Upper West	Upper West	Upper West
	District/ Municipalit y	Kumbung u	Tolon Savelug u		Bolgatang a	Kassena - Nankana	Bongo	Daffiama -Bussie- Issa	Lambussie -Karni	Nadowli -Kaleo
	Coordinates	9° 57'N 1° 02'W	-	9 <sup>°</sup> 59'N 0 <sup>°</sup> 85'W	10° 45'N 0° 50'W	10 <sup>°</sup> 52'N 1 <sup>°</sup> 08'W	10 <sup>°</sup> 52' N 0 <sup>°</sup> 51' W	10 27'N 02 34'W	10 <sup>°</sup> 40'N 02 <sup>°</sup> 38'W	10 <sup>°</sup> 11'N 02 <sup>°</sup> 36'W
Area of Wat	ershed (km <sup>2</sup> )	165	53	31	1.70	650	136	21	35	141

 Table 1: Description of Study Watersheds

Rainfall	Туре	Uni-modal	Uni-modal	Uni-modal
System	Annual Mean (mm)	1,000 – 1,300	700 – 1,010	800 - 1,100
	Duration (months)	5 - 6	5 - 6	5 - 6
Temperatur	Day	33 - 39	20 - 22	29.0
e ( <sup>©</sup> C)	Night	35 - 45	23 - 28	32.2
	Mean	33 - 45	23 - 30	30.0
Relative	Dry Season	50	10	20
Humidity (%)	Wet Season	80	65	70
Agro-ecologi	cal Zone	Guinea Savannah	Guinea/Sudan Savannah	Guinea Savannah
Geology		Precambrian basement rocks and Palaeozoic rocks from the Voltaian sedimentary basin	Metamorphic and igneous rocks with gneisss, graodiorite and sandstone	Precambrian, granite and metamorphic rocks
Soil Classes		Acrisols, plinthosols, planosols, luvisols, gleysols and fluvisols	Plinthosols, luvisols, vertisols, leptosols, lixisols, and fluvisols	Lixisols, fluvisols, leptosols, vertisols, acrisols and plinthosols

# 2.2 Integrated GIS-based RUSLE Model

The Integrated GIS-based RUSLE model was used to estimate the annual soil loss rate at the nine (9) watersheds. The model estimated annual soil loss rate by a cell-by-cell multiplication using raster maps of five (5) parameters as expressed in Equation 1 according to Renard*et al.* (1997). The model was simulated using ArcGIS with the detail process of the methodology illustrated in Figure 2. Input data included soil and rainfall, digital elevation model and landsat images of landuse/landcover of the watersheds.

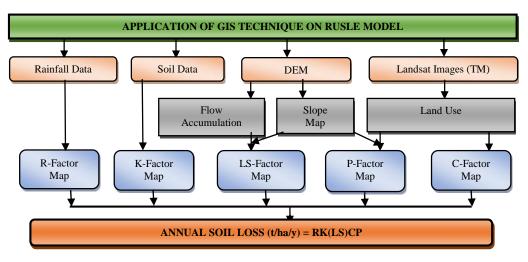
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## A = RK(LS)CP

## (1)

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Where: A - Annual soil loss (t/ha/y), R - Rainfall erosivity factor (MJmm/h/ha/y), K - Soil erodibility factor (t/ha/h/MJ/mm), LS - Slope length and steepness factor (dimensionless), C - Land cover management factor (dimensionless, ranges from 0 to 1), and P - Soil conservation support practice factor (dimensionless, ranges from 0 to 1).



**Figure 2:**Methodological Framework of Soil Loss Estimation in Watersheds Using RUSLE Model Adapted fromGelagay and Minale (2016)

## 2.2.1 Rainfall Erosivity Factor (R)

The R-factor quantifies the effect of rainfall intensity impact on soil erosion (Wischmeier and Smith, 1978) and it is estimated using Equation 2 (Hurni, 1985).

# R = 0.562(Ar) - 8.12

Where: R - Rainfall erosivity factor (MJmm/h/ha/y) and Ar - Annual rainfall (mm).

Twenty (20) years annual rainfall data obtained from the Ghana Meteorological Agencies rain gauges was interpolated by inverse distance weighted method to generate uninterrupted rainfall data for each 30 m grid cell in ArcGIS 10.4 environment. From this continuous rainfall data, the R-factor value of each grid cell was computed using raster calculator geo-processing tool.

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## 2.2.2 Soil Erodibility Factor (K)

The K factor represents soil type to erosion susceptibility, sediment transportability and amount and rate of runoff given a particular rainfall input (Ganasri and Ramesh, 2015). The soil maps of the watersheds were produced using Ghana soil shapefile (HWSD, 2017) with the region of interest extracted using each watershed area and the soil data layer clipped onto the watershed in ArcGIS environment10.4. The soil classes for each watershed were then determined from the soil maps and K-factor values obtained from literature for each soil class and then input into ArcGIS and processed to obtain K-factor maps for each watershed. The K-factor values (Table 2) for the corresponding soil classes were used in the study.

Table 2:Soil Erodibility (K) Factor Values for Different Soil Classes

FAO Soil Class	K-factor Value
Acrisols	0.25
Fluvisols	0.30
Leptosols	0.28
Lixisols	0.23
Planosols	0.34
Plinthosols	0.26
Vertisols	0.15

Adapted from Ashiagbor et al. (2014)

## 2.2.3 Slope Length and Steepness (LS) Factor

The slope length and steepness (LS) factor represents soil loss due to combinations of slope length and steepness relative to a standard unit plot (Ashiagbor*et al.*, 2014). The Spatial Analyst Extension Toolbox (Surface) in ArcGIS 10.4 was used to derive the slopes of the watersheds from the Digital Elevation Model (DEM) of the watershedsat 30 m resolution. All sinks in the DEM were identified and filled with the filled DEM for each watershed used as input to determine the flow direction (FD) and used as an input grid to derive the flow accumulation (FA). The LS-factor was calculated using raster calculator in ArcGIS and Equation 3 (Wischmeir and Smith, 1978).

$$LS = \left[ \left( \frac{Q_a M}{22.13} \right)^3 X \left( 0.065 + 0.045 S_g \right) + \left( 0.0065 S_g^2 \right) \right]$$

Where: LS - Slope length and steepness factor (dimensionless), Qa - Flow accumulation grid, M - Grid size, Sg - Grid slope (%), y - A dimensionless exponent (0.2 – 0.5) with varying values for different slopes depending on the slope steepness, being 0.5 for slopes exceeding 4.5 %, 0.4 for 3 - 4.5 % slopes, 0.3 for 1 – 3 %, and 0.2 for slopes less than 1 %.

## 2.2.4 Land Cover Management (C) Factor

The C factor quantifies the combined effect of plants, crop sequence and other soil cover surface on soil erosion (Molla and Sisheber, 2017). The C-factor is dimensionless with values ranging from 0 to 1. The C-factor maps were quantified from the landuse/landcover classes of the watersheds whilst the C-factor values (Table 3) by Hurni (1985) were used in this study.

#### Table 3:Land Cover Management (P) Factor

Landuse/Landcover Type	C-factor Value
Cropland	0.27
Built-up Areas	0.25
Water Body	0.00
Closed Savannah Woodland	0.05
Open Savannah Woodland	0.15

Adapted from Hurni (1985)

## 2.2.5 Erosion Management Practice (P) Factor

The study adopted a combination of general landuse and landcover types and slope. The P-factor values were assigned by categorizing the watersheds into major kinds of landuse and landcover types and assigned considering local management practices together with values in Table 4 (Sharma and Goyal, 2013) to produce the P-factor maps of the various watersheds. The P-factor is dimensionless with values ranging from 0 to 1.

(3)

Landuse/Landcover Type	P-factor Value
Cropland	0.50
Built-up Areas	0.80
Water Body	0.00
Closed Savannah Woodland	1.00
Open Savannah Woodland	1.00

**Table 4:** Erosion Management Practice (P) Factor

Adapted from Sharma and Goyal (2013)

#### 3. Results and Discussion

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# 3.1 The RUSLE Model Factors of the Watersheds

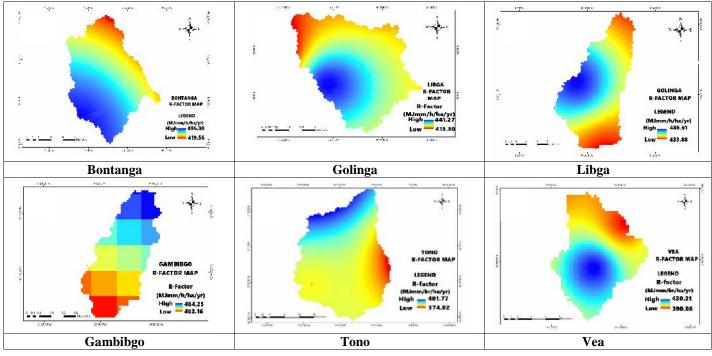
The estimated RUSLE model parameters which are R, K, LS and P factors of the nine (9) watersheds are presented in Table 5.

<b>Table 5:</b> Estimated RUSLE Factors of the Study Watersheds.
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Watershed	Mean Annual	<b>R-Factor Value</b>	K-Factor Value	LS-Factor	C-Factor	P-Factor
	Rainfall (mm)	(MJmm/h/ha/y)	(t/ha/h/MJ/mm)	Value	Value	Value
Bontanga	1074	419.56 - 486.08	0.25 - 0.34	0 - 2.16	0 - 0.27	0 - 1.0
Golinga	1030	433.88 - 486.61	0.25 - 0.34	0 - 1.50	0 - 0.27	0 - 1.0
Libga	1065	413.97 - 441.27	0.25 - 0.34	0 - 1.75	0 - 0.27	0 - 1.0
Gambibgo	969	403.16 - 404.25	0.23 - 0.30	0 - 1.24	0 - 0.27	0 - 1.0
Tono	970	374.92 - 401.77	0.15 - 0.30	0 - 6.36	0 - 0.27	0 - 1.0
Vea	934	390.86 - 430.31	0.15 - 0.30	0 - 2.57	0 - 0.27	0 - 1.0
Daffiama	1046	374.14 - 417.46	0.15 - 0.28	0 - 0.84	0 - 0.27	0 - 1.0
Karni	1042	379.28 - 428.93	0.15 - 0.28	0 - 2.07	0 - 0.27	0 - 1.0
Sankana	1009	373.61 - 447.05	0.15 - 0.28	0 - 1.73	0 - 0.27	0 - 1.0

#### 3.2 Rainfall Erosivity (R) Factor

The mean annual rainfall amounts in the 9 watersheds ranged from 934 mm to 1074 mm.Dabral*et al.* (2008) andGanasri and Ramesh (2015) noted that soil loss rate in the watersheds is more sensitive to rainfall. An estimated low R-factor value of 373.61 MJmm/h/ha/y at Sankana watershed and 486.61 MJmm/h/ha/y was estimated at the Golinga watershed as the highest. Details of the R-factor values are as presented in Table 5. Farhan *et al.* (2013) reported the distribution of R-factor values to vary and consistent with annual rainfall, and classified R-factor values of 300 – 600 MJmm/h/ha/y as moderately erosive. Rainfall in the watersheds is moderately erosive and might cause moderate soil loss. As shown in Figure 3, the influence of rainfall erosivity factor on soil loss in the watersheds decreases from the deep blue area to the deep red area.



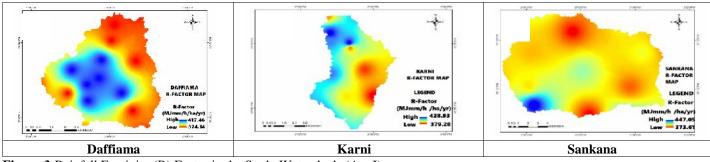


Figure 3: Rainfall Erosivity (R) Factor in the Study Watersheds (A - I)

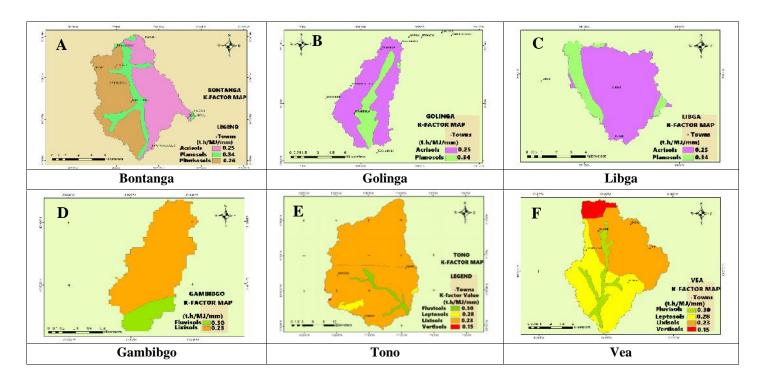
# 3.3 Soil Erodibility (K) Factor

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The number of soil classes identified in the study watersheds varied from 2 to 4 with the details presented in Table 6 and Figure 4. The K-factor values for the various soils ranged from 0.15 - 0.34 t/ha/h/MJ/mm with the lowest being vertisols and the highest being planosols. Based on the classification of NRCS-USDA (2002), the K-factor values obtained indicate that the watersheds have low to moderate erodible soils. The areal coverage statistics presented in Table 6 indicated that low erodible soils constituted the largest area of the Tono, Vea, Gambibgo, Libga and Golinga watersheds whereas the moderately erodible soils constituted the largest area of the Bontanga, Sankana, Karni and Daffiama watersheds This suggests that the contribution of K-factor to soil loss in the watersheds ranged from low to moderate.

Table 6: Soil Classes, K-factor Values, Erodility Class and Areal Coverage in the Study Watersheds

FAO Soil	K-factor	Erodibility	Name of Watershed/FAO Soil Class Areal Coverage (km <sup>2</sup> )										
Class	(t/ha/h/M J/mm)	Class	Bontanga	Golinga	Libga	Gambibgo	Tono	Vea	Daffiama	Karni	Sankana		
Lixisols	0.23	Low	-	-	-	1.40	610.30	60.60	-	0.14	47.30		
Fluvisols	0.30	Moderate	-	-	-	0.30	23.20	13.10	-	-	-		
Leptosols	0.28	Moderate	-	-	-	-	16.00	53.50	16.00	27.68	73.47		
Vertisols	0.15	Low	-	-	-	-	0.50	8.80	5.00	7.44	20.23		
Acrisols	0.25	Low	68.00	40.60	25.00	-	-	-	-	-	-		
Plinthosols	0.26	Moderate	69.50	-	-	-	-	-	-	-	-		
Planonosls	0.34	Moderate	27.50	12.40	6.00	-	-	-	-	-	-		
TOTAL	-	-	165.0	53.0	31.0	1.70	650.0	136.0	21.0	35.0	141.0		



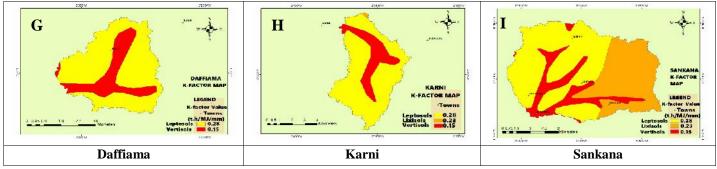


Figure 4: Soil Erodibility (K) Factor in the Study Watersheds (A – I)

## 3.4 Slope Length and Steepness (LS) Factor

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As presented in Table 5 and Figure 5, the combined LS-factor value of the watersheds varies from 0 to 6.36 (dimensionless) with the highest value estimated at the Tono watershed due to high terrains in some parts of the watershed. The estimated LS-factor values in all the watersheds indicate that the topography of the watersheds is relatively flat to gentle. According to Molla and Sisheber (2017), relatively flat to gentle topography has low LS-factor values ranging from 0 - 10.0 and does not contribute significantly to soil loss. This suggests that the influence of the combined slope length–steepness to soil loss might be low across all the study watersheds.

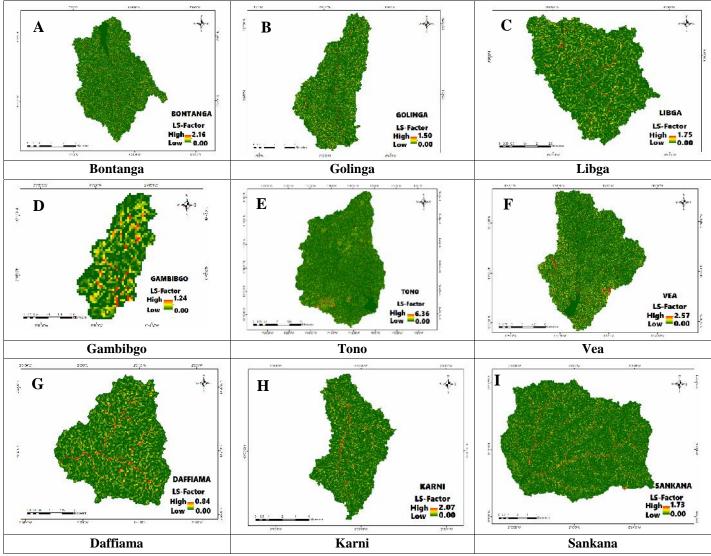


Figure 5:Slope Length and Steepness (LS) Factor in the Study Watersheds (A - I)

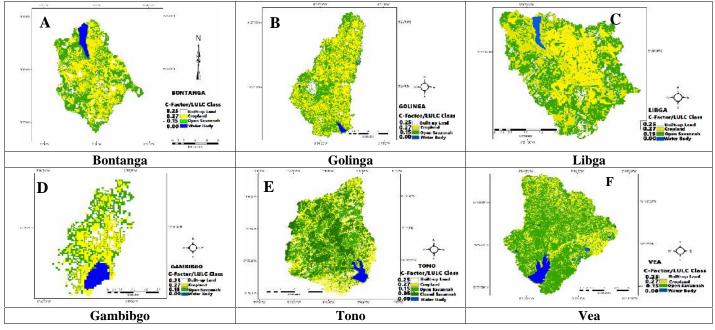
## 3.5 Landcover Management (C) Factor

Four (4) major landuse/landcover (LULC) classes thus; cropland, water body, built-up land and open savannah woodland were identified (Table7) in the study watersheds except Tono which had closed savannah woodland as the fifth major LULC class. The C-factor values of the LULC classes ranged for 0.0 - 0.27 and according to Renard*et al.* (1997), LULC classes with C-factor values above 0.20 significantly contributes to soil loss if conservation measures are not installed. Across all the watersheds, the predominant LULC was cropland with the highest C-factor value (0.27) and constituted over 40 % of the area of each watersheds (Table 7 and Figure 6). Cropland can significantly influence high soil loss in watersheds as tilling of the land for crop production destroys the vegetative cover as well as disturbs the soil and renders it susceptible to erosion and subsequent loss. Open and closed savannah woodlands with high potentials of controlling soil erosion and loss are noted to be depleted annually by crop production and built-up areas. Molla and Sisheber (2017) noted the contribution of C-factor to soil loss in watersheds to be higher in cultivated lands followed by built-up/bare lands and grasslands that are heavily grazed.

 Table 7: Different Landuse/Landcover Classes and their C-Factor and Areal Coverage

LULC	C-		Name of Watershed/LULC Areal Extent (km <sup>2</sup> )										
	Factor	Bontanga	Golinga	Libga	Gambibgo	Tono	Vea	Daffiama	Karni	Sankana			
CL	0.27	111.00	39.46	17.95	0.70	245.86	74.86	10.58	22.25	91.26			
BL	0.25	8.23	3.77	4.04	0.63	79.43	10.38	3.23	3.01	4.40			
WB	0.00	7.41	0.40	0.23	0.11	16.19	5.59	0.11	0.20	0.38			
OSW	0.15	38.36	9.37	8.78	0.26	198.60	45.17	7.08	9.54	44.20			
CSW	0.05	-	-	-	-	109.92	-	-	-	-			
Total	-	165.0	53.0	31.0	1.70	650.0	136.0	21.0	35.0	141.0			
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LULC – Landuse/landcover; CL – Cropland; BL – Built-up lan;, WB – Water bodies; OSW – Open savannah woodland; CSW – Closed savannah woodlan;, C-Factor – Landcover management factor



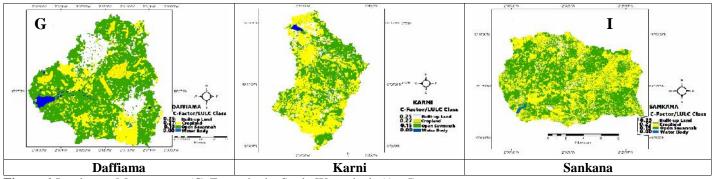


Figure 6:Landcover Management (C) Factor in the Study Watersheds (A - I)

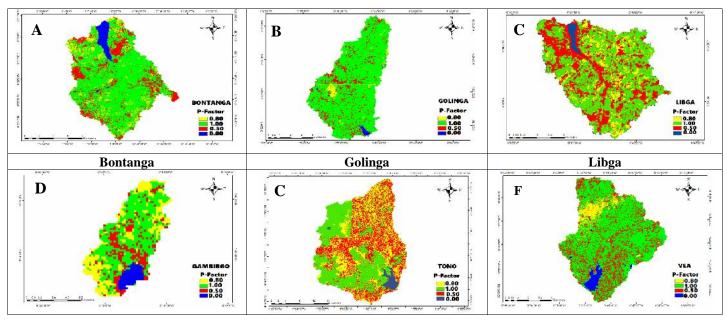
## 3.6 Erosion Management Practice (P) Factor

The P-factor values of the study watersheds ranged from 0 - 1.0 with water body recording the lowest value while the closed and open savannah woodlands recorded the highest value (Table 8 and Figure 7). According to Ganasri and Ramesh (2015), a P-factor value approaching zero indicates good soil conservation practice whereas the value approaching 1.0 indicates poor soil conservation practice in reducing or controlling soil loss due to erosion. The P-factor values obtained showed that no conservation practices are carried out to control soil loss in the closed and open savannah woodlands. However, some few soil conservation management practices like ploughing across the slope and earthen/stone bunding were being carried out in some croplands and built-up areas to check soil erosion and loss. Onori*et al.* (2006) indicated that cultivated and built-up lands with some soil conservation practices have P-factor values less than 1.0, while those without soil conservation management practices equals to 1.0.

 Table 8: Different Landuse/Landcover Classes and their P-Factor and Areal Coverage

LULC	P-Factor			]	Name of Watershed/LULC Areal Extent (km <sup>2</sup> )									
		Bontanga	Golinga	Libga	Gambibgo	Tono	Vea	Daffiama	Karni	Sankana				
CL	0.50	111.00	39.46	17.95	0.70	245.86	74.86	10.58	22.25	91.26				
BL	0.80	8.23	3.77	4.04	0.63	79.43	10.38	3.23	3.01	4.40				
WB	0.00	7.41	0.40	0.23	0.11	16.19	5.59	0.11	0.20	0.38				
OSW/	1.00	38.36	9.37	8.78	0.26	308.52	45.17	7.08	9.54	44.20				
CSW														
Total	-	165.0	53.0	31.0	1.70	650.0	136.0	21.0	35.0	141.0				

LULC – Landuse/landcover; CL – Cropland; BL – Built-up land; WB – Water bodies; OSW – Open savannah woodland; CSW – Closed savannah woodland; C-Factor – Landcover management factor



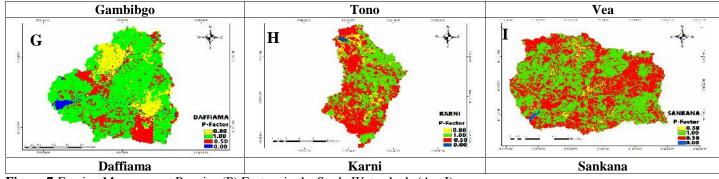


Figure 7: Erosion Management Practice (P) Factors in the Study Watersheds (A - I)

#### 3.7 Annual Soil Loss in Watersheds of Upper East Region

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The soil loss maps of the watersheds are shown in Figure 8 while the statistical details of the annual soil loss and their associated severity classes are presented in Table 9. The study estimated the annual soil loss at the Tono watershed to range from 0.0 - 96.30 t/ha/y with the mean loss as 6.91 t/ha/y. The annual soil loss at Vea watershed was estimated to range from 0.0 - 85.10 t/ha/y withmean 8.17 t/ha/y while the loss at the Gambibgo watershed ranged from 0.0 - 66.20 t/ha/y with mean of 5.37 t/ha/y. The estimated mean annual soil loss at the Vea watershed exceeded the FAO tolerable soil loss limits of 4.20 - 7.20 t/ha/y for soils with deep depth (FAO, 1984), but lower than the global tolerable soil loss limit of 11.20 t/ha/y (USDA-NRCS, 1999). However, the mean annual soil loss estimated at the Tono and Gambibgo watersheds were within the FAO and global tolerable soil loss limits. Abubakari (2014) estimated a very high mean annual soil loss of 22.83 t/ha/y at the Tono watershed using GeoWEPP model. Atakora*et al.* (2013) estimated mean annual soil loss of 6.8 - 10.2 t/ha/y at the watershed of Biemso valley in Ghana and this relatively conforms to the findings of this study. In a similar study in the Densu Basin located in the south-eastern part of Ghana, Owusu (2012) estimated mean annual soil loss of 2.20 t/ha/y, which is quite lower than the values in this study. This could be attributed to differences in agro-climatic zones.

As presented in Table 8, the severity of annual soil loss in the watersheds ranged from very low (0.0 - 1.0 t/ha/y) to very high (> 60.0 t/ha/y) with the dominant one being the very low class with a coverage area of 50.6 - 67.5 %. It, however, was noted to contribute the least of about 7.80 - 9.40 % of the total annual soil loss. The moderate severity soil loss class with a coverage area between 1.70 - 18.80 % was noted as the highest contributor of about 25.2 - 51.2 % of the total annual soil loss at the Tono and Gambibgo watersheds. At Vea watershed, the very high severity soil loss class contributed the highest of 29.1 % to its total annual soil loss (Table 9), and this could be due to the presence of leptosols and fluvisols which are less resistant to erosion. It was also observed from the study that the hot spot areas which contribute high to very high soil losses in the watersheds were spatially found in the steep slope parts of the watersheds and along the banks of water courses such as rivers and streams. Farhan *et al.* (2013) for example reported a strong correlation between the highest soil loss values and slope steepness in Kufranja watershed in Northern Jordan.

Wate	rshed		Gam	ıbibgo			Т	0 <b>no</b>		Vea					
Soil Loss (t/ha/y)	Soil Loss Severity Class	Coverage Area (ha)	% Total Area	Annual Soil Loss (t)	% Total Annual Soil Loss	Coverage Area (ha)	% Total Area	Annual Soil Loss (t)	% Total Annual Soil Loss	Coverage Area (ha)	% Total Area	Annual Soil Loss (t)	% Total Annual Soil Loss		
0.0 - 1.0	Very Low	86.0	50.6	77.4	9.4	39,913.3	61.4	37,917.6	8.5	9,173.9	67.5	8,715.2	7.8		
1.1 - 2.0	Low	35.2	20.7	104.0	10.4	3,124.1	4.8	24,635.3	5.5	1,471.5	10.8	17,828.4	16.1		
	Moderate	26.8	15.8	208.6	22.9	6,720.6	10.4	72,419.3	16.2	868.6	6.4	20,208.6	18.2		
2.1 - 5.0	Low														
5.1 - 25.0	Moderate	13.1	7.7	229.8	25.2	12,232.3	18.8	229,359.6	51.2	234.1	1.7	17,693.5	15.9		
25.1 - 60.	High	7.3	4.3	196.8	21.6	2,238.9	3.4	40,818.3	9.1	1,788.3	13.1	32,343.5	29.1		
> 60.0	Very High	1.7	1.0	95.9	10.5	770.8	1.2	43,043.6	9.6	63.6	0.5	14,246.3	12.8		
Total	-	170	100	912.5	100	65,000	100	448,193.8	100	13,600	100	111,053.5	100		
Mean	-	-	-	5.37	-	-	-	6.91	-	-	-	8.17	-		

Table 9: Severity and Coverage of Annual Soil Loss in Watersheds of Upper East Region

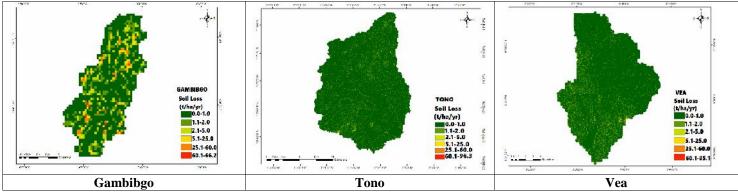


Figure 8:Spatial Distribution of Soil Loss in Watersheds of Upper East Region

#### 3.8 Annual Soil Loss in Watersheds of Upper West Region

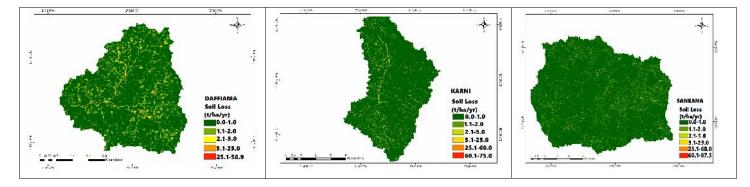
The annual soil loss rates at the Sankana, Karni and Daffiama watersheds which are situated in the Upper West Region of Ghana were estimated to range from 0.0 - 87.50 t/ha/y with mean loss ranging from 5.94 - 7.78 t/ha/y as presented in Figure 9. The estimated mean annual soil loss at Sankana and Daffiama watersheds were within the FAO (1984) soil loss tolerable limits of 4.20 - 7.20 t/ha/y and also lower than the global soil loss tolerable limit of 11.20 t/ha/y as reported by USDA-NRCS (1999). The mean annual soil loss at the Karni watershed was observed to have exceeded the FAO soil loss tolerable limits but lower than that of the global tolerable limit.

In a similar study in the Densu Basin in Ghana, Owusu (2012) estimated mean annual soil loss of 2.20 t/ha/y, which is quite lower than the estimated means in this study. Also, Mesele (2015) estimated lower mean annual soil loss of 1.42 t/ha/y at Kumasi-Anwomaso in Ghana. The very wide differences in means could be attributed to different geographical locations as the Densu Basin and Kumasi-Anwomaso are respectively located in the south-eastern and southern parts of Ghana whilst Daffiama, Karni and Sankana watersheds are located in the Guinea agro-ecological zone in northern Ghana.

The study also found very low to very high severity soil losses in all the watersheds except Daffiama. Across all the watersheds, the very low soil loss severity class constituted the largest area of about 36.70 - 48.30 % but contributes the least of about 5.9 - 7.9 % of the total annual soil loss in the watersheds (Table 10). Again, the moderate soil loss severity class was found contributing the highest between 31.2 - 38.6 % of the total annual soil loss of the watersheds.

Wate	rshed		Daf	fiama			Ka	rni		Sankana				
Soil Loss (t/ha/y)	Soil Loss Severity Class	Coverage Area (ha)	% Total Area	Annual Soil Loss (t)	% Total Annual Soil Loss	Coverage Area (ha)	% Total Area	Annual Soil Loss (t)	% Total Annual Soil Loss	Coverage Area (ha)	% Total Area	Annual Soil Loss (t)	% Total Annual Soil Loss	
0.0 - 1.0	Very Low	770.7	36.7	732.2	5.9	1,690.5	48.3	1,605.9	7.9	6,105.3	43.3	5,800.0	6.5	
1.1 - 2.0	Low	606.9	28.9	2,421.5	19.4	857.5	24.5	4,661.8	17.1	4,173.6	29.6	15,297.1	17.3	
	Moderate	212.1	10.1	2,227.1	17.9	381.1	10.9	5,340.8	19.6	1,762.5	12.5	17,799.3	20.1	
2.1 - 5.0	Low													
5.1 - 25.0	Moderate	310.8	14.8	4,817.4	38.6	430.5	12.3	9,631.8	35.4	1,015.2	7.2	27,690.4	31.2	
25.1 - 60.0	High	199.5	9.5	2,274.3	18.2	133.0	3.8	3,614.3	13.3	860.1	6.1	12,623.2	14.3	
> 60.0	Very High	-	-	-	-	7.0	0.2	2,350.1	8.6	183.3	1.3	9,416.8	10.6	
Total	-	2,100	100	12,472.5	100	3,500	100	16,714.6	100	14,100	100	88,626.8	100	
Mean	-	-	-	5.94	-	-	-	7.78	-	-	-	6.29	-	

Table 10:Severity and Coverage of Annual Soil Loss in Watersheds of Upper West Region



Daffiama	Karni	Sankana						
Figure 9: Spatial Distribution of Soil Loss in Watersheds of Upper West Region								

#### 3.9 Annual Soil Loss in Watersheds of Northern Region

As presented in Figure 10, the estimated annual soil loss rate ranged from 0.0 - 70.0 t/ha/y in the Bontanga, Golinga and Libga watersheds which are located in the Northern Region. The estimated mean annual soil loss at the watersheds of 3.71 - 4.96 t/ha/y presented in Table 11 were within the tolerable soil loss limits of 4.20 - 7.20 t/ha/y for soils with deep depth reported by FAO (1984) and also within the global tolerable soil loss limit of 11.2 t/ha/y indicated by USDA-NRCS (1999). However, it was observed that the estimated mean annual soil loss rates at the study watersheds were quite higher than the mean soil loss of 2.20 t/ha/y estimated by Owusu (2013) at the Densu Basin in located the south-eastern part of Ghana.

The severity of soil loss in the watersheds ranges from very low (0.0 - 1.0 t/ha/y) to very high (> 60.0 t/ha/y). The very low severity class was predominant with a coverage area ranging from 42.40 – 60.20 % and contributing about 7.70 – 15.4 % of the total annual soil loss (Table 11). The moderate severity class which constituted an area ranging from 3.50 - 15.20 % was noted as the highest contributor (25.0 - 42.80 %) of the total annual soil loss. As seen in Table 11, the watersheds were not much affected by very high soil loss as only a small area ranging from 0.10 - 1.10 % was affected by this soil loss severity class, contributing averagely 5.41 % of the total annual soil loss. This could be attributed to the relatively flat topography of the watersheds as well as the dominance of acrisols and plinthosols which are group 'A' soils and intrinsically less susceptible to severe erosion (USDA-NRCS, 1985). Farhan *et al.* (2013) stated that soil loss in watersheds varies spatially with changes in topography, soil characteristics and landuse/landcover.

Wate	Bontanga			Golinga					Libga					
Soil Loss (t/ha/y)	Soil Loss Severity Class	Coverage Area (ha)	% Total Area	Annual Soil Loss (t)	% Total Annual Soil Loss	Coverage Area (ha)	% Total Area	Annual Soil Loss (t)	% Total Annual Soil Loss		Coverage Area (ha)	% Total Area	Annual Soil Loss (t)	% Total Annual Soil Loss
0.0 - 1.0	Very Low	6,996	42.4	6,296.4	9.7	2,899.1	54.7	2,724.2	11.0		1866.2	60.2	7,747.5	15.4
1.1 - 2.0	Low	3,646.5	22.1	10,903.0	13.3	1,245.5	23.5	2,478.5	13.8		911.4	29.4	2,588.7	22.8
	Moderate	2,970.0	18.0	17,285.4	21.1	535.3	10.1	4,073.6	16.4		173.6	5.6	2,233.3	19.7
2.1 - 5.0	Low													
5.1 - 25.0	Moderate	2,508.0	15.2	35,011.7	42.8	455.8	8.6	10,241.8	37.4		108.5	3.5	2,832.2	25.0
25.1 - 60.0	High	264.0	1.6	6,937.9	8.5	106	2.0	3,209.7	12.9		37.2	1.2	1,827.1	16.1
> 60.0	Very High	115.5	0.7	5,325.7	6.5	58.3	1.1	2,157.1	8.7		3.2	0.1	117.8	1.04
Total	-	16,500	100	81,760.1	100	5,300	100	24,885.9	100		3,100	100	11,346.6	100
Mean	-	-	-	4.96	-	-	-	4.70	-		-	-	3.71	-

Table 11: Severity and Coverage of Annual Soil Loss in Watersheds of Northern Region



Figure 10:Spatial Distribution of Soil Loss in Watersheds of Northern Region

#### 4. Conclusions

The study used the RUSLE model integrated with remote sensing and GIS tools to estimate the magnitude of annual soil loss and the spatial distribution of soil loss severity classes in nine (9) watersheds in northern Ghana. The estimated annual soil loss in the various watersheds ranged from 0 - 96.30 t/ha/y. Estimated mean annual soil loss of 7 watersheds were within tolerable soil loss limits according to FAO (1984) and USDA-NRCS (1999) classification. Mean annual soil loss estimated for 2 watersheds i.e. Vea and Karni exceeded the tolerable soil loss limit set by the USDA-NRCS (1999). The severity of annual soil loss rates in the watersheds ranged from very low class (0.0 - 1.0 t/ha/y) to very high class (> 60.0 t/ha/y). Across all the watersheds, the very low severity class constituted the largest area of 36.70 - 67.50 % but was noted as the least contributor of 5.90 - 15.40 % to the total annual soil loss. At Tono and Vea watersheds, which have leptosols and fluvisols noted to be less resistant to soil erosion as well as the high terrain in some parts watersheds presented high and very high severity soil loss classes and contributed significantly to the total annual soil loss in the watersheds.

Farming practices such as ploughing along the slope, slashing and burning, and farming very close to the banks of water courses in the watersheds were some of the main causes of soil loss in the watersheds. High terrain and slope steepness in some parts of Tono, Vea watersheds also contributed to high soil loss. Soil and water conservation measures such as contour ploughing, stone/earth bunding and upstream afforestation of the watersheds are best practices to reduce soil loss. Education on riparian area protection and avoidance of farming in buffer zones of the reservoirs which has a multiplier negative effect on soil loss and reservoir sedimentation is very necessary in all watersheds. Analysis of the temporal distribution of the soil loss in the watersheds is necessary and therefore recommended for future studies.

#### References

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- Abbott, M. B., Bathurst, J. C., Cunge, J. A., O'Connell, P. E. and Rasmussen, J. 1986. An introduction to the European Hydrological System-Systeme Hydrologique Europeen 'SHE' 1: History and philosophy of a physically-based, distributed modelling system. *Journal of Hydrology*, Vol. 87, pp. 45-59.
- Abubakari, A. 2014. Impact of Land Use Changes on Soil Erosion and Sedimentation in the Tono Reservoir Watershed using GeoWEPP Model. Master of Science Thesis, Kwame Nkrumah University of Science and Technology.
- Amegashie, B. K. 2009. "Assessment of Watershed Erosion, Sedimentation and Nutrient Export intoSmall Reservoirs from their Watersheds in the Upper East Region of Ghana. A Master of Science Thesis, Kwame Nkrumah University of Science and Technology, Kumasi, Ghana. Accessed 20<sup>th</sup> December, 2017.
- Arnold, J. G., Allen, P. M. and Bernhardt, G. 1993. A comprehensive surface-groundwater flow model. *Journal of Hydrology*, Vol. 142, No. 4, pp. 47-69.

Ashiagbor, S., Forkuo, E. K., Laari, P. and Aabeyir, R. 2014. Modelling soil erosion in the densu river basin using RUSLE and GIS Tools. *Journal of Environment, Science and Engineering*, Vol. 56, No. 3, pp. 247-254.

- Atakora, E. T., Kyei-Baffour, N., Ofori, E. and Antwi, B. O. 2013. "simulation of sediment transport to sawah rice fields by applying the water erosion prediction project (WEPP) model to a watershed in Ghana. *Journal of Soil Science and Environmental Management*, Vol. 4, No. 3, pp. 46-53.
- Bai, Z. G., Dent, D. L., Olsson, L., and Schaepman, M. E. 2008. Proxy global assessment of land degradation. Soil Use and Management, 24, 223–234.
- Beck, M. B. 1987. "Water quality modeling: a review of the analysis of uncertainty. *Water Resources Research*, Vol. 23, No. 8, 1393-1442.
- Bennett, J. P. 1974. Concepts of mathematical modeling of sediment yield. *Water Resources Research*, Vol. 10, No. 3, pp. 485-492.

Dabral, P. P., Baithuri, N., and Pandey, A. 2008. "Soil Erosion Assessment in a Hilly Watershed of North Eastern India using USLE, GIS and Remote Sensing". *Water Resources Management*, 22(12), 1783-1798.

- FAO, 1984. "Ethiopian Highland Reclamation Study". Final Report, Vol 1-2. Rome.
- Farhan, Y., Zregat, D., and Farhan, I. 2013. Spatial estimation of soil erosion risk using RUSLE approach, RS, and GIS Techniques: A case study of Kufranja Watershed, Northern Jordan. *Journal of Water Resource and Protection*, Vol. 5, No. 12, p. 1247.
- Ganasri, B. P. and Ramesh, H. 2015. Assessment of soil erosion by RUSLE model using remote sensing and GIS A case study of Nethravathi Basin, *Geoscience Frontiers xxx*, pp 1 9, <u>http://dx.doi.org/10.1016/j.gsf.2015.10.007</u>
- Gelagay, H. S. and Minale, A. S. 2016. Soil loss estimation using gis and remote sensing techniques: A case of Koga watershed, Northwestern Ethiopia". *International Soil and Water Conservation Research* 4, pp 126–136, http://dx.doi.org/10.1016/j.iswcr.2016.01.002
- Hajigholizadeh, M., Melesse, A. and Fuentes, H. 2018. Erosion and sediment transport modelling in shallow waters: A review on approaches, models and applications. *International Journal of Environmental Research and Public Health*, Vol. 15, No. 3, pp. 518: 1-24
- Hurni, H., 1985. Erosion-productivity-conservation systems in Ethiopia. In Proceedings of Paper Presented at the 4<sup>th</sup> International Conference on Soil Conservation, 3-9 November 1985. Maracay, Venezuela.
- HWSD, 2017. Harmonised World Soil Databse (HWSD version 1.21).
- Kandel, D. D., Western, A. W., Grayson, R. B. and Turral, H. N. 2004. Process parameterization and temporal scaling in surface runoff and erosion modelling. *Hydrol. Process.* Vol. 18, No. 8, pp. 1423-1446.
- Kothyari, U. C., Tewari, A. K., Singh, R. 1994. Prediction of sediment yield. *Journal of Irrigation and Drainage Engineering* ASCE Vol. 120, No. 6, pp. 1122-1131.
- Laflen, J. M., Lane, L. J., and Foster, G. R. 1991. WEPP: A new generation of erosion prediction technology. Journal of Soil and

Water Conservation, Vol. 46, No. 1, pp. 34-38.

- Lal, R. 1994. Soil Erosion Research Methods; CRC Press: Boca Raton, FL, USA, 1994.
- Mekonnen, M., Keesstra, S. D., Baartman, J. E., Stroosnijder, L. and Maroulis, J. 2017. Reducing sediment connectivity through man-made and natural sediment sinks in the minizr watershed, Northwest Ethiopia. *Land Degradation & Development*, Vol. 28, No. 2, pp. 708-717.
- Mesele, S. A. 2015. Quantifying maize yield and erosion influencing factors for soil loss prediction under different tillage and soil amendments. Master of Science Thesis Submitted to the Kwame Nkrumah University of Science and Technology, Kumasi, Ghana.
- Molla, T. and Sisheber, B., 2017. Estimating Soil Erosion Risk and Evaluating Erosion Control Measures. Copernicus Publications, the European Geosciences Union, Solid Earth, pp 14-25, doi: 10.5194/se-8-13-2017.
- Morgan, R. P. C., Quinton, J. N., Smith, R. E. Govers, G., Poesen, J.W. A., Chisci, G. and Torri, D. 1998. The EUROSEM Model. In Modelling Soil Erosion by Water; Springer: Berlin/Heidelberg, Germany, pp. 389-398.
- Nearing, M. A., Jetten, V., Baffaut, C., Cerdan, O., Couturier, A., Hernandez, M., Le Bissonnais, Y., Nichols, M. H., Nunes, J. P., Renschler, C. S., Souche're, V., van Oost, K. 2005. Modeling response of soil erosion and runoff to changes in precipitation and cover. *Catena*, Vol. 61, pp. 131-154.
- Onori, F., Bonis, P. D. and Grauso, S. 2006. Soil erosion prediction at the basin scale using the revised universal soil loss equation (RUSLE) in a watershed of Sicily (southern Italy). *Environ. Geol.*, 50(8), 1129-1140.
- Owusu, G. 2012. A GIS-based estimation of soil loss in the Densu Basin in Ghana. *West African Journal of Applied Ecology*, Vol. 20, No. 2, pp. 41-51.
- Renard, K. G., Foster, G. R., Weesies, G. A., McCool, D. K. and Yoder, D. C. 1997. "Predicting Soil Erosion by Water: A Guide to Conservation Planning with the Revised Universal Soil Loss equation", Agriculture Handbook No. 703, U.S. Department of Agriculture, Washington, D.C., USA.
- Schulze, R. E. 1995. Hydrology and Agrohydrology: A Text to Accompany the ACRU 3.00 Agrohydrological Modelling System; Water Research Commission: Pretoria, South African".
- Sharma, G. and Goyal, R. 2013. Qualitative and Quantitative Soil Erosion Mapping of Micro-watersheds of Bisalpur Reservoir using Remote Sensing and GIS. UIM-2013.
- Shiferaw, A. 2011. Estimating soil loss rates for soil conservation planning in the BorenaWoreda of South Wollo highlands, Ethiopia. *Journal of Sustainable Development in Africa*, Vol. 13, No. 3, pp. 87-106.
- Shinde, V., Tiwari, K. N. and Singh, M. 2010. Prioritization of Micro Watersheds on the Basis of Soil Erosion Hazard using Remote Sensing and Geographic Information System. *International Journal of Water Resources and Environmental Engineering*, Vol. 5, No. 2, pp. 130-136.
- United States Department of Agriculture Natural Resources Conservation Service (USDA-NRCS), 1999. National Soil Survey Handbook: title 430-VI. US Government Printing Office, Washington DC.
- USDA-NRCS, 1985. "Hydrology". National Engineering Handbook, Supplement A, Section 4, Natural Resource Conservation Service, USDA, Washington, D.C.
- USDA-NRCS, 2002. Technical Guide to RUSLE Use in Michigan. Institute of Water Research, Michigan State University. Accessed 10/08/2017 at <u>http://www.iwr.msu.edu/rusle/kfactor.htm</u>
- Wang, H. W., Kondolf, M., Tullos, D., and Kuo, W. C. 2018. Sediment management in Taiwan's reservoirs and barriers to implementation. *Water*, Vol. 10, No. 8, p. 1034.
- Wheater, H. S., Jakeman, A. J. and Beven, K. J. 1993. Progress and Directions in Rainfall-Runoff Modelling; FAO: Rome, Italy.
- Wischmeier, W. H. and Smith, D. D. 1978. Predicting Rainfall Erosion Losses A Guide to Conservation Planning; Agriculture Handbook No. 537; US Department of Agriculture Science and Education Administration: Washington, DC, USA, 1978; p. 168.

#### **Biographical notes**

T. A. Adongo is of the Department of Agricultural Mechanisation & Irrigation Technology, University for Development Studies, Tamale, Ghana

W. A. Agyare and N. Kyei-Baffour are of the Department of Agricultural and Biosystems Engineering, Kwame Nkrumah University of Science and Technology, Kumasi, Ghana

F. K. Abagale is of the Department of Agricultural Engineering, University for Development Studies, Tamale, Ghana

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